

Radioactive Decay Worksheet

Alpha decay: nucleus spontaneously emits an alpha particle (symbol: α particle), which is 2 p⁺ and 2 n (or also the same as a Helium (He) atom).

Result: atomic number decreases by 2 (lost 2 p⁺)

Result: atomic mass decreases by 4 (lost 2 p⁺ and 2n = 4 amu)

Beta decay: neutron in nucleus spontaneously emits a beta particle (symbol: β particle), which is essentially an electron trapped in a neutron. The neutron, therefore, turns itself into a proton.

Result: atomic number increases by 1 (gained 1 p⁺)

Result: atomic mass stays same (no mass lost or gained: β particle or electrons have no mass)

Beta or electron capture: proton in nucleus captures a beta particle (symbol: β particle), which is essentially an electron that can become part of a neutron. The proton, therefore, turns itself into a neutron.

Result: atomic number decreases by 1 (lost 1 p⁺)

Result: atomic mass stays same (no mass lost or gained: β particle or electrons have no mass)

Example

| Original | alpha decay | beta decay | alpha decay | beta capture | beta decay | alpha decay |
|----------|-------------|------------|-------------|--------------|------------|-------------|
| 85 | 83 | 84 | 82 | 81 | 82 | 80 |
| At | Bi | Po | Pb | Tl | Pb | Hg |
| Astatine | Bismuth | Polonium | Lead | Thallium | Lead | Mercury |
| 210 | 206 | 206 | 202 | 202 | 202 | 198 |

Complete this table

| Original | beta decay | alpha decay | beta capture | alpha decay | alpha decay | beta decay |
|----------|------------|-------------|--------------|-------------|-------------|------------|
| 90 | | | | | | |
| Th | | | | | | |
| Thorium | | | | | | |
| 232 | | | | | | |

Complete this table

| Original | beta capture | alpha decay | alpha decay | beta capture | alpha decay | beta decay |
|----------|--------------|-------------|-------------|--------------|-------------|------------|
| 92 | | | | | | |
| U | | | | | | |
| Uranium | | | | | | |
| 238 | | | | | | |

Radiometric Dating Worksheet

When radioactive isotopes (parent - P) decay, they produce daughter products (D) at a constant rate, called the half-life (T). Example: if we start with 100 atoms of the parent, after one half-life, there will be 50 parent atoms remaining and 50 daughter atoms newly made. After another half-life (two half-lives), there will be 25 parent atoms remaining and now 75 daughter atoms. Each parent-daughter isotope pair has its own half-life. To achieve the above example with U-238 takes 9 billion years (two half-lives). To achieve the above example with C-14 takes 11400 years (two half-lives). In the geologic environment, we use a mass spectrometer to count the number of Parent and Daughter atoms in a closed-system (like minerals crystallizing from magmas), and use the relative proportions to find out how old the closed-system is.

- Assuming we start with only parent isotopes (no daughter), after one half-life has passed, there should be $\frac{1}{2}$ parent remaining and $\frac{1}{2}$ daughter newly formed. The ratio of P:D is $\frac{1}{2} : \frac{1}{2}$ or 1:1. Complete the rest of this table, as in the first example:

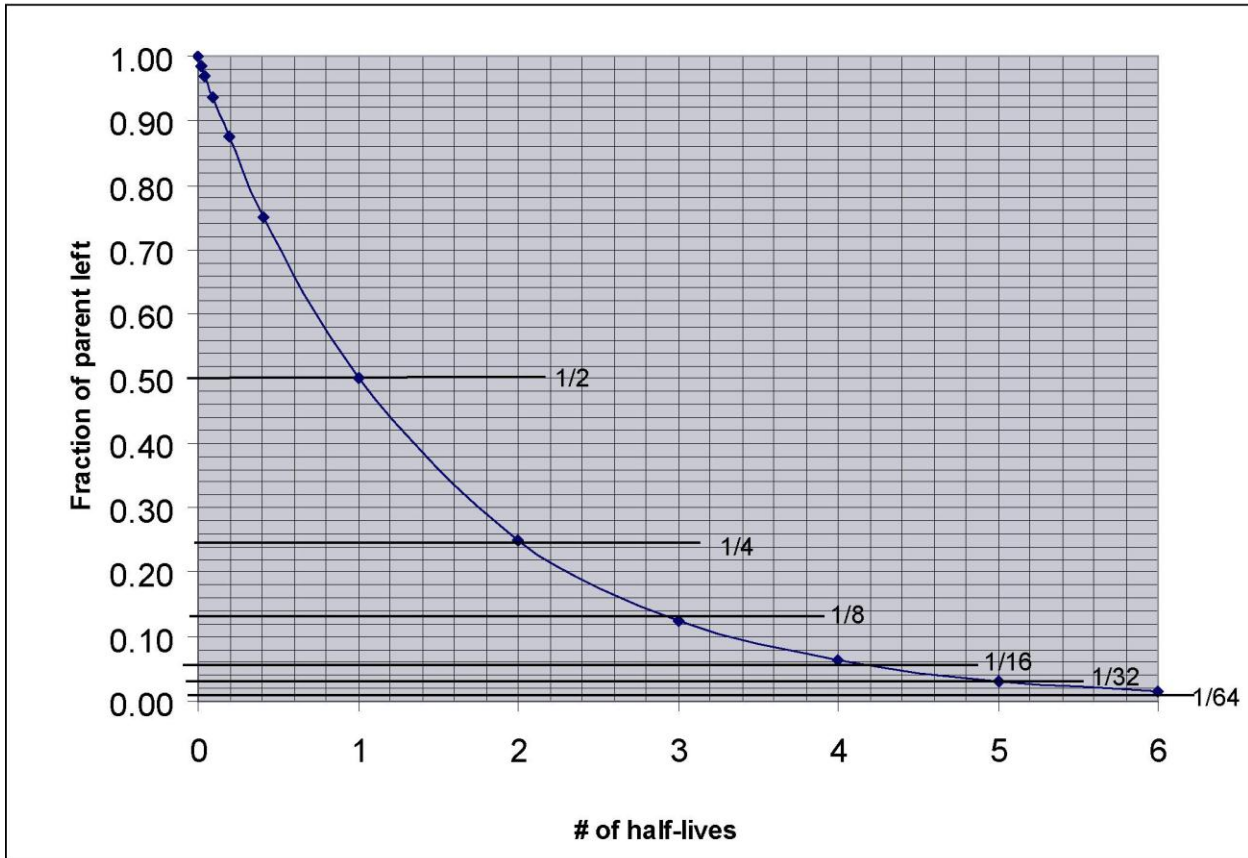
| # Halflives | Fraction of original Parent remaining | Fraction of original parent turned into daughter | Parent:Daughter ratio |
|-------------|---------------------------------------|--|-----------------------|
| 1 | $\frac{1}{2}$ | $\frac{1}{2}$ | 1:1 |
| 2 | | | |
| 3 | | | |
| 4 | | | |
| 5 | | | |
| 6 | | | |

Not all rocks can be dated radiometrically. Some because they cannot maintain closed systems (like metamorphic rocks); others because they do not contain radioactive isotopes (like quartz sandstones); and finally some because the radioactive isotopes that they do contain have half-lives that are either too long or too short to be measured for a rock of a certain age (like trying to date a 1 m.y.-old rock by using C-14 decay - which would have been completely decayed after about 150,000 years).

| Parent (P) | Daughter (D) | Half-lives ($T_{1/2}$) | Materials dated |
|------------|--------------|--------------------------|--|
| U-238 | Pb-206 | 4.5×10^9 yr | Zircon (igneous rocks - source; and sedimentary rocks as grains) |
| U-235 | Pb-207 | 0.7×10^9 yr | Zircon (igneous rocks - source; and sedimentary rocks as grains) |
| K-40 | Ar-40 | 1.4×10^9 yr | Micas, volcanic rock (igneous rocks) |
| C-14 | N-14 | 5700 yr | Shells, limestone, organic materials |

- To date the age of a shell found in an old Indian fishing village, which isotope pair would you measure? Why?
- If you want to date a meteorite, which isotope pair would you measure? Why?
- If you want to date lava flows on an old lava flow on Kauai (probably about 8 m.y.), which isotope pair would you measure? Why?
- If you want to date zircon crystals in ancient sandstones in Australia, which isotope pair would you measure? Why?

6. If the C-14:N-14 ratio in a shell in a sandstone was found to be 1:3, how old is the shell?
7. If the U-235:Pb-207 ratio in a zircon in a sandstone was found to be 1:3, how old is the zircon?
8. If the K-40:Ar-40 ratio in a zircon in a granite was found to be 1:1, how old is the sample?
9. If the U-238:Pb-206 ratio in a zircon in a lava flow was found to be 3:1, how old is the flow?



| T (# of Halflives) | Fraction Parent | Daughter | Ratio |
|--------------------|-----------------|----------|-------------|
| 0 | 1 | 0 | infinity :1 |
| 0.0227 | 63/64 | 1/64 | 63:1 |
| 0.0458 | 31/32 | 1/32 | 31:1 |
| 0.0931 | 15/16 | 1/16 | 15:1 |
| 0.1927 | 7/8 | 1/8 | 7:1 |
| 0.4151 | 3/4 | 1/4 | 3:1 |
| 1.0000 | 1/2 | 1/2 | 1:1 |
| 2.0000 | 1/4 | 3/4 | 1:3 |
| 3.0000 | 1/8 | 7/8 | 1:7 |
| 4.0000 | 1/16 | 15/16 | 1:15 |
| 5.0000 | 1/32 | 31/32 | 1:31 |
| 6.0000 | 1/64 | 63/64 | 1:63 |

CURVE EQUATION: $T = -1.443 \ln(f)$ f = fraction of parent left; T = # of half lives that have passed